

## MINERALOGY OF STARDUST COMETARY GRAIN T112: SIMILARITY TO AMOEBOID OLIVINE AGGREGATES.

KOMATSU, Mutsumi<sup>1\*</sup>, MIKOUCHI, Takashi<sup>2</sup>, FAGAN, Timothy<sup>1</sup>, ZOLENSKY, Michael<sup>3</sup>

<sup>1</sup>Waseda University, <sup>2</sup>Department of Earth and Planetary Science, Graduate School of Science, University of Tokyo, <sup>3</sup>NASA Johnson Space Center

### Introduction:

The successful analysis of comet 81P/Wild 2 particles returned by the Stardust mission has revealed that the Wild 2 dust contains abundant silicate grains that are much larger than interstellar grains and appear to have formed in the inner regions of the solar nebula [1]. Wild 2 particles include minerals which are isotopically and mineralogically similar to CAIs [e.g., 2], chondrules [e.g., 3], and amoeboid olivine aggregates (AOAs) [e.g., 4] in chondrites. In this study, we have examined the mineralogy of the Track 112 particle and compared the possible relationships between T112 and AOAs.

### Results and Discussion:

Our two TEM grids from T112 are dominated by 4 micron -sized forsteritic olivine. Submicron sized chromite is associated with the olivine. The olivine from T112 is near end-member forsterite, but shows a slight enrichment of Fe toward the grain rim. In the least metamorphosed chondrites, most AOAs consist of forsterite. Enrichment of olivine in fayalite component is correlated with petrologic subtype of a host meteorite [5]. In CV chondrites, Fe-enrichment is distinct in Allende (petrologic type >3.6). The Fe enrichment on the rim of T112 olivine may have been caused by a minor degree of thermal processing after the condensation of forsterite.

From EDS analyses, the chromite grains have high Cr<sub>2</sub>O<sub>3</sub> content (up to 65 wt. %), and yield Al-absent compositions with intermediate Mg# ; however, it is likely that some Mg detected by EDS is from neighboring olivine and that the grains are closer to pure chromite in composition. Chromite is a minor phase in Wild 2 particles. Cr-rich spinels have been identified in the chondrule-like particle Torajiro [3] and associated with Coki-B Kool fragments (T141 [6]); however, these spinels have significant Al<sub>2</sub>O<sub>3</sub>.

Three possibilities for the formation of chromite can be considered: (i) direct condensation from a gas; (ii) crystallization from chondrule melt; (iii) metamorphism/aqueous alteration. (i) Modeling by [7] predicts that Cr-rich spinel crystallizes in mixed vapor + silicate liquid + refractory solid reactions at high temperatures in dust enriched systems, but always with significant Al<sub>2</sub>O<sub>3</sub>. (ii) This modeling [7] also shows that chromite is stable with silicate liquid at high dust enrichment, but textures of T112 particle do not appear to be typical of chondrules. Furthermore, the 16O-rich composition of olivine in the potted butt sample from (T112) is more consistent with AOAs than chondrules. (iii) Tiny Cr-rich grains surrounding olivine in type 3 ordinary chondrites were identified by [8] as products of incipient metamorphism. The T112 chromite might have a similar origin.

The forsteritic, 16O-rich composition of T112 olivine suggests a condensation origin. Slight enrichment in Fa-content along the olivine rim suggests some metamorphism. At this point, we consider metamorphism the most likely origin of T112 chromite.

References: [1] Brownlee D. E. et al. (2006) *Science* 314:1711-1716. [2] Zolensky M. et al. (2006) *Science* 314:1735-1739. [3] Nakamura T. et al. (2008) *Science* 321:1664-1667. [4] Nakamura-Messenger K. et al. (2011) *MaPS* 46:1033-1051. [5] Komatsu M. et al. (2001) *MaPs* 36:629-641. [6] Joswiak D. J. et al. (2009) *MaPS* 44: 1561-1588. [7] Ebel D. and Grossman L. (2000) *GCA* 64:5339-366. [8] Grossman J. N. and Brearley A. J. (2005) *MaPS* 40:87-122.

Keywords: meteorites, carbonaceous chondrites, cometary dust