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Speciation and partitioning of C-bearing volatiles species in and between silicate-saturated C-O-H fluids and (C-O-H)-saturated melts have been determined in-situ with the samples to pressures and temperatures of ~2GPa and 900°C, respectively. Structural characterization was conducted with vibrational spectroscopy of samples contained in externally-heated, hydrothermal diamond anvil cells. The redox conditions were controlled near that of the Fe+H<sub>2</sub>O=FeO+H<sub>2</sub> (reducing, RED) and Ni+H<sub>2</sub>O=NiO+H<sub>2</sub> (oxidizing, OX) equilibria, respectively. Melts are, therefore saturated in H<sub>2</sub>O, H<sub>2</sub>, and C-bearing species (redox dependent) and coexisting fluids saturated in silicate components. Solubility of volatile and silicate components depend on both temperature and pressure.

The melt/fluid partition coefficients of the C-bearing species vary with redox conditions and temperature with the  $\Delta H_{RED}^{melt/fluid} = 44(7)$  kJ/mol and  $\Delta H_{OX}^{melt/fluid} = -70(32)$  kJ/mol. Pressure is a dependent variable and increases with increasing temperature. It is assumed no pressure effect of the partition coefficients.

The solution equilibria under reducing and oxidizing conditions, respectively, were; (1)  $2CH_3^- + H_2O + Q^{n+1} = 2CH_4 + Q^n$  and (2)  $2CO_3^{2-} + H_2O + 2Q^{n+1} = HCO_3^- + 2Q^n$ , where the superscript, n, in the Q-species denotes number of bridging oxygen in the silicate species (Q-species). In the absence of H<sub>2</sub>O equilibrium (1) changes to  $CH_3^- + Q^n = CH_4 + Q^{n+1}$ . For oxidized carbon, there is an analogous expression expressing equilibrium between molecular CO<sub>2</sub> and structurally bound CO<sub>3</sub><sup>2-</sup>-groups. Under both oxidizing and reducing conditions, the abundance ratios, CH<sub>4</sub>/CH<sub>3</sub><sup>-</sup> and HCO<sub>3</sub><sup>-</sup>/CO<sub>3</sub><sup>2-</sup> increase with temperature. The enthalpy change associated with the species transformation does, however, differ for fluids and melts and also for oxidized and reduced carbon ( $\Delta H_{(1)}^{fluid} = -16(5)$  kJ/mol,  $\Delta H_{(1)}^{melt} = -49(5)$  kJ/mol,  $\Delta H_{(2)}^{fluid} = 81(14)$  kJ/mol). For the exchange equilibrium of CH<sub>4</sub> and CH<sub>3</sub><sup>-</sup> species, the temperature-dependent equilibrium constant yields  $\Delta H = 34(3)$  kJ/mol.

Reactions (1) and (2) involve changes in silicate polymerization where increasing abundance ratios, CH<sub>4</sub>/CH<sub>3</sub><sup>-</sup> and CO<sub>3</sub><sup>2-</sup>/HCO<sub>3</sub><sup>-</sup> lead to increased silicate melt polymerization. As a result of the relations between speciation of C-bearing species and melt and fluid structure, stable isotope (C and H) and element partition coefficients between melts and fluids, which depend on and silicate polymerization and silicate speciation, also vary with speciation of C-bearing species in silicate-C-O-H systems. Pressure, temperature, and redox control on the C-speciation also govern those (and other) properties.

キーワード: COH volatiles, fluid structure, melt structure, high pressure, high temperature, redox conditions  
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