

The role of fluid for diamond-free UHP carbonate and calc-silicate rocks from the Kokchetav Massif

Kazumasa Aoki[1]; Yasuhide Inoue[1]; Minoru Kikuchi[1]; Yoshihide Ogasawara[1]

[1] Earth Sci., Waseda Univ.

H₂O-rich fluid played an important role for the formation of diamond-free rocks under UHP conditions. Local heterogeneity of fluid compositions during UHP metamorphism has been well demonstrated in cm-scaled sample of diamond-free banded marble. Diamond-free garnet-clinopyroxene rocks that look like a skarn also suggested very low XCO₂ under UHP conditions.

The diamond-free banded marble (no. Y665) consists of three zones; zone-A: Ti-clinohumite-bearing dolomitic marble, zone-B: dolomite marble, zone-C: dolomite marble that lacks Ti-clinohumite and forsterite. All three zones lack diamond. Zone-A is characterized by the presence of Ti-clinohumite and by Ti-clinohumite-aragonite tie-line at the peak metamorphic conditions. Diopside lacks any lamellae. Large-grained Ti-clinohumite shows clear compositional zoning of TiO₂ between the core (1.5-1.7 wt.%) and the rim (2.0-3.3 wt.%). Zone-B indicated the stability of dolomite-diopside pair. No lamella was observed in diopside. Zone-C contains diopside with phlogopite and garnet lamellae. These lamellae were confirmed with a laser Raman spectroscopy and an electron microprobe. A retrograde product, high MgCO₃ calcite with large-grained exsolved dolomite lamellae indicating precursor aragonite + dolomite assemblage, occurs in zone-C. This means dolomite-diopside pair was stable under UHP conditions in zone-C. All three zones have different bulk chemical compositions each other; however, the tie-line relations are unable to be explained only by their bulk compositions. In zone-C and B, diopside-dolomite pair was stable, whereas an incompatible assemblage aragonite-Ti-clinohumite was stable in zone-A. Aragonite-clinohumite is stable at very low XCO₂, and diopside-dolomite has wide stability range in XCO₂ that is higher than that of aragonite-clinohumite. The difference in the distribution of Ti-bearing phases in each zone was detected by chemical mappings. Many of Ti-clinohumite grains occur in zone-A (low XCO₂ side); however, a few grains of Ti-bearing phases in zone-B and C (relatively high XCO₂ side). Such paragenetic relations and the distribution of Ti-bearing phases in banded specimen (no. Y665) strongly suggest the heterogeneous fluid compositions and the fluid origin of TiO₂ during UHP metamorphism. The banded structure of this Kokchetav marble might be created by such kind of fluid effect.

Garnet-clinopyroxene rocks (no. XX16) consists mainly of garnet, clinopyroxene, quartz, amphibole, K-feldspar, and symplectite after clinopyroxene with minor amounts of titanite, calcite. No diamond occurs in this calc-silicate rock. Coesite exsolution needles and plates were found in titanite in this sample. Coesite was identified with a laser Raman spectroscopy. The mode of occurrence of the coesite in titanite is similar to that reported in the calcite marble from Kumdy-Kol; however, this garnet-pyroxene rock is clearly different from the calcite marble in modal composition and the chemical compositions of clinopyroxene; this indicates that protoliths for those rocks were different each other. Clinopyroxene contains K-feldspar and phengite lamellae that were identified by the laser Raman spectroscopy and the electron microprobe.

Summarizing all these data described above, diamond-free marble and calc-silicate rocks had been subjected to UHP metamorphism with strong effect of H₂O-rich fluid. The H₂O-rich fluid could be initiated by the dehydrations in surrounding gneisses and eclogites in the Kumdy-Kol area.