

## Is a high $^3\text{He}/^4\text{He}$ derived from the depleted mantle? - constraint from kimberlites

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MORBs (mid-atlantic oceanic basalts) and OIBs (oceanic island basalts) have systematically different noble gas isotope ratios and they show  $^3\text{He}/^4\text{He}$  ratios of about  $8R_a$  ( $1R_a$ : the  $^3\text{He}/^4\text{He}$  ratio of the atmospheric value) and  $5\text{--}40R_a$ , respectively. Since alpha-decay of U and Th produces radiogenic  $^4\text{He}$ , old continental crusts enriched in U and Th show  $^3\text{He}/^4\text{He}$  ratios lower than those of MORBs. Low  $^3\text{He}/^4\text{He}$  ratios observed in OIBs are regarded to have been affected by recycled slabs into the mantle. However, high  $^3\text{He}/^4\text{He}$  ratios observed in OIBs should represent a magma source with an averaged  $^3\text{He}/(\text{U}+(\text{Th}))$  higher than that of MORBs through geologic time. In that case, there are two possibilities about the source that the primordial  $^3\text{He}$  remained relatively abundant or depleted in U and Th. In the latter case, however, there should be a phase where U is more depleted than He. Since OIBs are more abundant than MORBs in U, it is generally regarded that the OIB magma source is less degassed.

On the other hand, if a reservoir is resided in a convective mantle, it seems hard to remain primordial noble gases in it. Hence, those who prefer a model of whole mantle convection to the other models have a serious problem to explain the occurrence of a reservoir where  $^3\text{He}$  remains. To explain it, it has been argued that a depleted phase with low U would retain  $^3\text{He}$  and ancient olivines composed of the mantle is a candidate. If noble gases were not lost during olivine formation in magma at an ancient time, olivines kept higher  $^3\text{He}/\text{U}$  ratios than that of the original source due to their partition coefficients and show a high  $^3\text{He}/^4\text{He}$  ratio at present. In such a case, however, the amount of He remained in olivines is less than 1% and even 0.01% of that of the original source. Since the magma source of OIBs is regarded to be located in the lower mantle, if the  $^3\text{He}$  content is low in the original magma source, it seems difficult to keep a high  $^3\text{He}/^4\text{He}$  ratio due to contamination at a shallow depth during a rise of a magma or plume. Furthermore, the MORB source shows more depleted character of solid element isotope signatures (e.g.,  $^{87}\text{Sr}/^{86}\text{Sr}$ ,  $^{143}\text{Nd}/^{144}\text{Nd}$ , etc.) than the OIB source.

It has been reported that the magma source of kimberlites has a  $^3\text{He}/^4\text{He}$  ratio of more than  $26R_a$  and the solid element isotopes cluster around the value of Bulk Earth. Hence, it is not reasonable to assign a depleted character to the kimberlite magma source. Based on the Nd-Hf isotope systematics, it has been suggested that kimberlite magma source would have been affected by a component which is originated from a reservoir produced by ancient E-type MORBs of more than 2Ga located in a non-convective mantle. In such a reservoir, radiogenic  $^4\text{He}$  with the amount of  $10\text{E}(-3\text{--}4)\text{cc/g}$  has been produced. To show a  $^3\text{He}/^4\text{He}$  ratio of more than  $26R_a$  in a kimberlite magma affected by such a component, the  $^3\text{He}$  with the amount of about  $10\text{E}(-7)\text{cc}$  is required. Since the highest concentration of  $^3\text{He}$  observed in a terrestrial rock is about  $10\text{E}(-9\text{--}10)\text{cc/g}$  in popping rocks, the above estimate indicates that kimberlite magma source should retain abundant  $^3\text{He}$ . Thus, based on information of kimberlites, it is inferred that a reservoir with high  $^3\text{He}/^4\text{He}$  magma source is not depleted in incompatible elements and contain relatively abundant  $^3\text{He}$ . Hence, it should be less degassed and less fractionated part located in a deep non-convective mantle.