Phase relations in the eclogite-carbonate-chloride system at 7.0-16.5 GPa: implications to mantle fluids and diamond inclusions

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Recent studies of fluid inclusions in diamonds and other mantle minerals indicate that chloride can be important in the mantle fluids. Complex chloride-carbonate-water fluids have been detected in many diamonds worldwide; however nature and speciation of these fluids are poorly constrained. Melting phase relations have been determined in a model chloride-carbonate-bearing eclogite at 7.0, 10.5, and 16.5 GPa and 1200-1800oC using multianvil technique. Starting composition was simplified Mid-Ocean-Ridge basalt in the system SiO2-Al2O3-MgO-CaO. 4 wt.% CO2 was added in the form of MgCO3 and CaCO3 and 3.56 wt.% Cl was added as NaCl and KCl.

The typical silicate phase assemblage coexisting with partial melts at 7.0-10.5 GPa is garnet-clinopyroxene-kyanite-stishovite. At 16.5 GPa kyanite and clinopyroxene were not stable and corundum was observed with garnet and stishovite. Unknown Cl-free Al-rich phase (carbonate?) was detected at 16.5 GPa. It has an approximate formula, recalculated to one carbon atom (calculated from electron microprobe analyses deficit), Mg0.3Al3.0Si0.7O5CO3. KCl and NaCl crystals were stable in low temperature runs coexisiting with magnesite and Ca- and Mg-bearing carbnatite melt. Melting of chloride produces chloride-carbonate melt immiscible with carbonatite. High temperature runs at 1600-1800oC reveal homogenous Si-rich carbonate-chloride partial melts. Although melting sequence of chloride and carbonate was found to be similar with that in peridotite-carbonate-chloride system (Litasov and Ohtani, 2009) melt compositions were significantly different. For example, at 7 GPa and 1550oC melt composition was (wt.%): SiO2=50.3, Al2O3=8.5, MgO=3.7, CaO=10.6, Na2O=1.5, K2O=4.0, CO2=17.5, Cl=4.9, O(-)=0.9. This composition correspond to Si-rich endmember of silicate-carbonate-bearing sequence of fluid inclusions in diamonds at Si,Al (silicate) - Mg,Fe,Ca (carbonate) - Na,K (chloride) triangle (e.g. Klein-BenDavid et al., 2007).

Chlorides were stable 100-200oC above the solidus. This fact indicates that melt composition in chloride-carbonate peridotite can be effectively controlled by the presence of water. In the anhydrous environments a low-fraction of Cl- and alkali-poor Ca-Mg-rich carbonatite melts can be formed and migrate from the source, whereas under the hydrous conditions, chloride can be dissolved into water to form alkali-rich chloride-carbonate brine-like melts. At higher temperatures (1400-1600oC) two immiscible carbonatite and chloride-carbonate liquids may coexist. The composition of carbonatite and chloride-carbonate melts is consistent with that of melt inclusions in fibrous/cloudy diamonds worldwide. Typical trends of melt evolution upon cooling and formation the melts included in diamonds may involve formation of immiscible Si-rich carbonatite and Si-rich chloride-carbonate melts from homogenous Cl-bearing kimberlite or carbonatite. The results are also applicable for carbonate and chloride stability and dynamics during deep subduction.