

Chronology and sedimentation of hyper salinity lake(Meedee lake) in the eastern Mediterranean

Naoki Izumitani[1]; Takuya Sagawa[2]; Masafumi MURAYAMA[3]; Hirofumi Asahi[4]; Hiroshi Kitazato[5]; Yasuyuki Nakamura[6]; Masaaki Shirai[7]; Juichiro Ashi[7]; Hidekazu Tokuyama[7]

[1] Natural Science, Kochi Univ.; [2] SRFC, Ehime Univ.; [3] Marine Core, Kochi Univ.; [4] ORI, University of Tokyo; [5] IFREE, JAMSTEC; [6] Ocean Res. Inst., Univ. Tokyo; [7] ORI, Univ. Tokyo

In the Mediterranean Sea, an evaporite rock accumulated was formed by Messinian Salinity Crisis during 5.33 to 6Ma (this event that the Mediterranean Sea was in isolation from the Atlantic, and desiccation) (Krijgsman et al., 1999). Saline materials exude from the evaporite rock and form ultra high salinity brine lakes in the basin. In this brine lake, it is a unique environment that salinity contents more than 300 psu which is equal to about 10 times of the seawater and low dissolved O₂. Five brine lakes have ever been discovered in the eastern Mediterranean. Microbiological research has been well studied in these lakes. The main purpose of this study is to reveal the sedimentary environment of a newly discovered brine lake. From the survey of the research group in KH06-4 cruise, a reverse fault was formed by tectonic motion with subduction. This fault reaches the evaporite rock which was dissolved by pore water. High salinity pore water elevates to the seafloor along the reverse fault. As a result, a brine lake was formed at a depression in the seafloor.

A piston core was collected at the edge of this lake with observation of the ocean floor by using a Navigable Sampling System during the KH06-4 cruise (latitude: 34°27.02'N, longitude: 22°16.61'E, W.D.: 2920m, core length: 293.5cm). The core consists of a calcareous ooze. No sedimentary structure is observed by a visual description and X-ray CT scan. Alternation of the dark color band (yellowish orange) and the light color band (grayish white) was observed on a several cm scale. The boundary of each band changes gradually or sharply, indicating a different change in sedimentary environments. After cutting half of the whole core, non-destructive measurements were conducted (magnetic susceptibility / digital color; *L *a *b/ color image / X-ray CT scanning). After cubes were sampled from the core continuously, water content and sand fraction (more than 0.063mm) were calculated. Two planktonic foraminifera were picked up and measured for stable isotopes using mass spectrometry. Benthic foraminifera were picked up from about 200 individuals and Benthic Foraminifer Number (BFN)/g was calculated.

Age was determined by AMS¹⁴C age of planktonic foraminifer (*G. inflata*) and the d¹⁸O of planktonic foraminifer (*G. ruber*) correlation to the standard oxygen isotope curve from the Vostok ice core. This core covers about 50 to 220 kyr, and the average sedimentation rate is calculated as about 2.0cm/kyr. Deposition time of each band was calculated as an average of several thousand years.

About the sedimentary environment, the surface sediment inside of the brine lake shows a white color and outside shows a dark color from the observation of NSS. Pyrite is found only in the light color band. Average number of BFN/g in the dark color band is about 260 and diversity is high. Average value of BFN/g in the light color band is about 128 and diversity is low. In addition, two types of species which tolerate low oxygen (*Stainforthia complanata*) and tolerate high salinity (*Articulina tubulosa*) are frequently found in the light color band. As above results, the dark color band and light color band were deposited in oxidation and reduction environments, respectively.

From comparison of the values between d¹⁸O of *G. ruber* and L*, the light color band corresponds to the interglacial periods and the dark color band corresponds to the glacial period. The changes in the oxidation-reduction environment are caused by: (1) the core was recovered from the edge of the brine lake, (2) the surface of the brine lake was elevated by inflow of pore water which was limited by the strength of the bottom current corresponding to the period of glacial-interglacial in the eastern Mediterranean Sea. When color changes occurred sharply and *L values did not correspond with d¹⁸O, it might be a sudden change of the topography with tectonic motions or following injection of the pore fluid from the deep.