

SSS023-P21

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## Estimation of Complex Spectral Ratio of Surface and Borehole Seismometry and Numerical Tests

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I compared 4 estimation methods of complex spectral ratios from seismic vertical array records and the results of numerical tests of noise-added synthetic records by these methods. Phase difference spectra are useful to identify velocity and attenuation structure using the spectral ratios. Examined estimation methods are two least-square approaches ( $H_1$  and  $H_2$ ), a geometrical mean of  $H_1$  and  $H_2$  ( $H_3$ ), and geometrical mean of the Fourier spectra ( $H_G$ ). These methods are different ways of assumption of noise.

The two least-square approaches assume that one of the surface or borehole records ( $y(f)$  and  $x(f)$  in the frequency domain, respectively) includes noise and the other one is noise-free. The least-square solution of the observational equation that is assumed that surface record  $y(f)$  includes noise is  $H_1=C_{xy}(f)/S_{xx}(f)$ , where  $C_{xy}$  is the ensemble mean of the cross spectra of  $x(f)$  and  $y(f)$  and  $S_{xx}$  is the one of the power spectra of  $x(f)$ . The solution that borehole records includes noise is  $H_2=S_{yy}(f)/C_{yx}(f)$ , where  $S_{yy}$  is the ensemble mean of the power spectra of  $y(f)$ . Phase difference spectra of  $H_1$  and  $H_2$  are identical.

The effects of the noise are shown in the expected value. The noise is no effects on the expected value of  $C_{xy}$ . This causes that the phase differences are expected to be robust. The expected values of  $S_{xx}$  and  $S_{yy}$  are affected by the variance of the noise of  $x(f)$  and  $y(f)$ , respectively. This indicates that the  $H_1$  and  $H_2$  are different for the same  $x(f)$  and  $y(f)$ .

I applied these 4 methods to estimate the spectral ratios of the spectra of noise-added synthetic surface and borehole records. The noise-free surface and borehole synthetic records were calculated with the exact transfer function of the vertically incident plain wave in the homogeneous medium with the surface. I made noise-added synthetic records by adding 20 different white noise on the noise-free synthetic records.

Stacked spectral ratios were calculated by the 4 methods. Mean values of cross and power spectra of the 20 pairs of noise-added synthetics were calculated.  $H_1$ ,  $H_2$ , and  $H_3$  were calculated from the mean values.  $H_G$  were calculated with geometrical mean of the spectral ratios of the Fourier spectra  $y(f)/x(f)$ . The result of  $H_1$  indicates small error in the frequencies between peaks of the exact transfer function, whereas, at the peak frequency of the transfer function,  $H_1$  underestimates the peaks due to the noise in the borehole records.  $H_2$  indicates clear peaks, whereas it overestimate the valleys.  $H_3$  are the geometrical mean value that is robust, but indicates unclear peak compared with  $H_2$ .  $H_G$  are close to  $H_3$ . Phase difference spectra of  $H_1$  and  $H_2$  are very close to the exact transfer function.

Smoothed spectra of noise-free synthetic spectra are very similar to the stacked spectra.  $H_1$ ,  $H_2$  and  $H_3$  were calculated from smoothed cross and power spectra.  $H_G$  are the geometrical mean of the spectral ratios of the smoothed Fourier spectra  $x(f)$  and  $y(f)$ . Parzen window was used as a smoothing window. The spectra applied both of stacking and smoothing also results similar to the stacked spectra.

Keywords: seismic vertical array, complex spectral ratio, transfer function