

Hcp 鉄の高温高圧弾性特性 High-P,T elasticity of hcp iron

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Earth's inner core (329~364 GPa and 5000~6000 K) is known to be composed of hexagonal closed pack (hcp) structured solid Fe-Ni alloy and seismologically anisotropic. Thermoelasticity of hcp (ϵ) iron is therefore a key to interpreting seismological information of the inner core. Since experimental measurements are still technically impractical, theoretical approaches in particular ab initio density functional computation have substantial roles. There are two different ways to simulate high-P,T elastic constants (c_{ij}) of crystals. One is based on the lattice dynamics method + quasiharmonic approximation (Sha and Cohen, 2010a,b) and the other is based on the molecular dynamics method (Vocadlo et al., 2009; Martorell et al., 2013). The former and the latter basically fail to capture higher-order anharmonicity and low-temperature quantum effects, which would be substantial and marginal in subsolidus condition, respectively. Due to these problems, distinct differences can be seen in high-P,T c_{ij} and their temperature dependences calculated by these different approaches. In this study, we performed ab initio molecular dynamics simulations employing a supercell containing 96 Fe, which is 50% larger than in the previous study with 64 atoms (Vocadlo et al., 2009; Martorell et al., 2013), to check the previous results. Technical details for computing high-P,T c_{ij} are basically the same as in our previous studies (Ichikawa et al., 2014; Kawai and Tsuchiya, 2015). We will present temperature dependences of elastic wave velocities and their anisotropies at the Earth's inner core pressures over 300 GPa.

Kawai and Tsuchiya, Geophys. Res. Lett. (2015) under review; Ichikawa et al., J. Geophys. Res., 119, 240 (2014); Martorell et al., Science 342, 466 (2013); Sha and Cohen, Phys. Rev. B 81, 094105 (2010a); Sha and Cohen, Geophys. Res. Lett. 37, L10302 (2010b); Vocadlo et al., Earth Planet. Sci. Lett. 288, 534 (2009)

キーワード: 第一原理計算法, Hcp 鉄, 弾性特性, 地球内核

Keywords: Ab initio calculation method, Hcp iron, Elasticity, Earth's inner core

Fcc FeHx at core pressure Fcc FeHx at core pressure

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水素は、地球の中心核に含まれる軽元素の候補として最も重要な元素のひとつである。地球形成初期の核形成時に、水素は選択的に金属コアに分配されたと考えられている (e.g. Fukai, 1984; Okuchi et al., 1997)。しかし、鉄-水素系の相関係に関する研究は比較的低压力の領域に限られており (Sakamaki et al., 2009)、実際の地球核において鉄-水素合金がとりうる安定な結晶構造は明らかにされていない。これまで、自由エネルギーの計算から、鉄-水素合金の構造は圧力の増加に伴って二重六方細密充填構造 (dhcp 構造) から六方細密充填構造 (hcp 構造)、さらに面心立方構造 (fcc 構造) へと変化することが予測されているが (Isaev et al., 2007)、未だ実験によって実証されたことはない。そこで本研究では、レーザー加熱式ダイヤモンドアンビルセルによる高温高压発生と高輝度放射光施設 SPring-8 における X 線回折測定により、鉄-水素合金の dhcp もしくは hcp 構造から fcc 構造への相転移を確認した。実験の結果、約 60 GPa において dhcp 構造から hcp 構造に変化し、約 70 GPa において hcp 構造から fcc 構造に変化することが確かめられた。さらに、我々は fcc 構造の鉄-水素合金について 26-137 GPa で格子体積を取得した。その結果、約 70 GPa において圧力と格子体積の関係および非圧縮率に不連続な変化がみられた。これは、理論計算から予測されている鉄-水素合金の磁気的性質の変化で説明できる可能性がある。今回得られた結果から、地球核の温度圧力条件における鉄-水素合金の結晶構造は従来提唱されてきた dhcp 構造ではなく、fcc 構造である可能性がある。

キーワード: 地球核, 水素, fcc 構造

Keywords: core, hydrogen, fcc structure, high pressure, X-ray diffraction

地球核条件下にレーザー衝撃圧縮された Fe-Si 合金の音速 Sound velocities of liquid Fe-Si alloys at Earth's core pressures by laser-shock compression

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Sound velocity at Earth's core conditions are one of the most important physical properties in Earth science because it can be directly compared with the seismological Earth model (PREM: Preliminary Reference Earth Model) [1]. The composition of solid inner core is estimated from the comparison of the model [1] and the extrapolation of sound velocities as a function of density of iron and iron alloys obtained by the static compression experiment [2, 3]. Birch's law, a linear sound velocity-density relation [4], is used to extrapolate sound velocities to densities in the core condition. On the other hand, the composition of liquid outer core is estimated from the partitioning and solubility data in the inner core boundary condition for the composition of solid core. There has been some works for the sound velocity of iron on the Earth's core condition by dynamic techniques using explosive [5], gas gun [5, 6], and laser [7]. However, the previous dynamic compression experiments are not enough to reveal the core of Earth, giant planets [8], and super-Earth which is at core pressures over 800 GPa [9]. In this study, we measure the sound velocity and density of liquid iron alloys by shock-compression method using high-power laser at pressures corresponding to super-Earth core pressures.

We conducted shock-compression experiments using a High Intensity Plasma Experimental Research (HIPER) system at the GEKKO-XII laser irradiation facility [10] at the Institute of Laser Engineering, Osaka University. The samples are Fe-Si alloys (Fe₉₅Si₅, Fe₉₀Si₁₀, Fe₈₀Si₂₀ and Fe₆₆Si₃₄ in weight percent). The sound velocities and densities of shock-compressed Fe-Si alloys using the high-power laser were measured by x-ray radiography [7, 11, 12] at pressures up to 960 GPa. The linear relation between the sound velocity and the density for FeSi alloy well follows Birch's law [4] up to 960 GPa along the Hugoniot. The extrapolated sound velocity of FeSi alloy was about 40% faster than that of PREM at inner core boundary pressure. The outer core is composed of Fe-Si alloy with 5-13 wt.% Si assuming Si is only light element at the core. This Si content is consistent with the results of previous work by sound velocity measurement [13] and shock-compression experiment [14].

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Keywords: sound velocity, laser, shock wave, Fe-Si alloys, Earth's core, experiment

液体 Fe-Ni-S 合金の弾性波速度 Sound velocity of liquid Fe-Ni-S alloy at high pressure: Sulfur in the core?

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The liquid Earth's outer core is composed mainly of iron (Fe)-nickel (Ni) alloy. Birch (1952) first found that the core is less dense than pure iron based on comparison between seismological observations and experimental measurements of the density of solid iron. This is the so called "core density deficit" problem. He suggested the existence of lighter component(s) in the core, and hydrogen (H), carbon (C), nitrogen (N), oxygen (O), silicon (Si), and sulfur (S) have been identified as likely candidates from cosmochemical and geochemical arguments (e.g., Poirier et al., 1994). The density difference between the outer core and the pure iron has been estimated to be 5-10%, depending on the assumed outer core geotherm (e.g., Anderson and Isaak, 2002). The nature of light elements has remained one of the biggest enigmas for the more than half-century since the Birch's work (1952). To justify the kind and quantity of the light elements in the core, sound velocity measurements of liquid iron alloying with possible lighter elements are fundamental because they link directly to seismological observations. We have launched the project on the sound velocity measurements for liquid iron alloys at high pressure in externally-heated and laser-heated diamond-anvil cells (DAC). The sound velocity of liquid (Fe,Ni)₃S was measured via a high resolution inelastic X-ray scattering (IXS) measurements at BL35XU of the SPring-8 synchrotron facility, Japan (Baron et al, 2001). We successfully determined the sound velocity of liquid (Fe,Ni)₃S up to the pressure of 50 GPa, which corresponds to the center of Mars. With our newly obtained results, we discuss the possibility of sulfur in the liquid cores of Earth and Mars.

キーワード: 弾性波速度, 液体鉄合金, 硫黄, 外核, 火星核

Keywords: Sound velocity, liquid iron alloy, sulfur, outer core, Martian core

鉄合金液体の第一原理熱弾性計算による地球外核組成の制約 Earth's outer core composition constrained by ab initio thermoelasticities of liquid Fe alloys

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The density deficit of the Earth's outer core indicates substantial amounts of light elements (O, Si, S, C, and H) incorporated in the outer core (Birch, 1952; 1964). The chemistry and amount of the light elements have been strongly debated for over 60 years. Ab initio molecular dynamics (AIMD) simulations have been widely applied to investigate several properties of liquid Fe and Fe alloys (e.g., Alfe et al., 2002; 2007; Badro et al, 2014; Ichikawa et al, 2014). Badro et al. (2014) recently reported a likely compositional model being consistent with seismological data. However with applying empirical pressure corrections, the model suggests smaller amount of light elements to reproduce the ICB density jump. In our study, adopting the Ichikawa et al. (2014) technique we determined the equations of state (EoS) of the liquid Fe alloys by means of the AIMD method in the P, T condition widely covering the entire outer core condition without any pressure corrections. From the EoS, densities, adiabatic bulk moduli, and finally P-wave velocities were calculated and compared with the seismological data (PREM) (Dziewonski and Anderson, 1981). After examining alloy systems from binary to quaternary, we could find some optimized compositional models. However, these have almost comparable reproducibility to PREM, suggesting that other observables are required to make further constraints on the outer core composition. If considering the observed large ICB density jump additionally, Fe-Ni-Si-O and Fe-Ni-S-O compositions appear the most likely.

Keywords: Earth's outer core compositional models, Ab initio molecular dynamics simulations, Equation of state of liquid Fe alloys

外核最下部にとくに敏感な波を用いた速度構造の推定 Fine seismic velocity structure of the lowermost outer core determined using outer core sensitive phases

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The liquid outer core of the Earth is regarded as homogeneous and approximately in hydrostatic equilibrium [Stevenson, 1987]. However, a low-velocity layer appears at the base of the outer core (F-layer) [Souriau and Poupinet, 1991; Kennett et al., 1995]. This basal layer may exhibit hemispherical features, as reported by Yu et al. [2005], corresponding to the quasi-hemispherical pattern of the inner core [Tanaka and Hamaguchi, 1997; Wen and Niu, 2002]. The pattern of the inner core is suggested to reflect solidification and melting at the inner core boundary (ICB), which might cause an Fe-rich or Fe-poor layer in the F-layer [Gubbins et al., 2008; Alboussiere et al., 2010; Monnereau et al., 2010]. However, the seismic profile of the F-layer is poorly revealed because of the non-uniqueness of the profiles investigated using previous methods and the interdependence of the F-layer velocity and other seismic properties of the Earth. Thus, a better constrained F-layer velocity is required before discussing its composition.

In this study, we investigated the velocity profile of the F-layer using two new methods: frequency dispersion of the traveltimes of waves that graze or are diffracted at the ICB, and differential traveltimes between waves reflected from the boundary and those that turn above the boundary. The first approach is sensitive to velocity gradients in the layer, while the second is sensitive to velocity excesses or deficits relative to a reference model for the layer; neither approach is sensitive to inner core properties or its radius. We analyzed seismograms of South American earthquakes observed using the Hi-net array [Okada et al., 2004] and the J-array network [J-Array Group, 1993] in Japan. The area investigated in the study is beneath the eastern Pacific, which is placed on the quasi-western hemisphere of the inner core.

Our results show that V_p values in the F-layer are intermediate between those of AK135 and PREM, and that the vertical velocity gradient is larger than that of AK135. Nearly constant velocities in the F-layer are not suited to observations.

キーワード: 外核最下部 P 波速度, F 層, PKPbc 波走時の分散, PKPbc と PKiKP との走時差

Keywords: P-wave velocity in the lowermost outer core, F-layer, Traveltime dispersion of PKPbc, differential traveltimes between PKiKP and PKPbc

生命居住惑星とプレートテクトニクスの役割 Role of Plate Tectonics for Habitable Planet

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Plate tectonics plays a role of global material circulation from the surface of the earth to the bottom of mantle since ca.4.0Ga. This is a due function both to bear habitable planet and to evolve as a habitable planet.

In spite of empirical recognition of importance of plate tectonics, nobody succeeded to synthesize the role of plate tectonics. However, the author completed this task by using following data; (1) geophysical constraints for the mechanism of plate tectonics, (2) petrological and geochemical characteristics of lithosphere, and (3) geologic history of the Earth, specifically the structural and petrological remarks of rock components and dynamics recorded in orogenic belts over the world.

As the result, basic condition for the operation of plate tectonics and following 6 roles of plate tectonics were summarized. The reason why plate tectonics is operated on the Earth is because mid-oceanic ridge is hydrated to enable plate to subduct which is helped by the lubricant water-rich fluids on the bottom of lithosphere. Roles of plate tectonics are (1) Global material circulation of CO₂ and H₂O, (2) Role of tectonic erosion, (3) Production of nutrients-source rocks at subduction zone, (4) Driving force of Earth's magnetic field, (5) The buffer of Earth's system, and (6) The controller of thermal history of the Earth.

地球内部の東西半球構造とグローバルダイナミクス East-west hemispherical structures in the Earth and their implications for global dynamics

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Hemispherical structures have been found in the inner core (Tanaka and Hamaguchi, 1997; Waszek et al., 2011; and the references therein), the outer core (Tanaka and Hamaguchi, 1993; Yu et al., 2005), and the mantle (Iwamori and Nakamura, 2012). While seismic velocities characterize the core hemispherical structures, the mantle east-west hemispheres have been proposed based on geochemistry, rather than south-north division as has been long argued for (Hart, 1984, known as “Dupal anomaly”). In order to better characterize and interpret the mantle geochemical hemispheres in both spatial and compositional domains, and to discuss whether the hemispherical structures in the core and mantle have any dynamical linkage or not, a total of 6854 young basalt data consisting of five isotopic ratios of Sr, Nd and Pb from almost all tectonic settings (mid-ocean ridge, ocean island, arc and continent) have been statistically analyzed (Iwamori and Nakamura, 2015).

As a result, it has been found that the continental basalts are mostly distributed only in the eastern hemisphere, while other basalts are distributed evenly. Using multivariate analysis (Independent Component Analysis, ICA), two independent compositional vectors have been extracted, which explain most of the sample variance (95%). Therefore, almost all young basalts from various tectonic settings plot on a single isotopic compositional plane, and can be explained solely by two elemental differentiation processes (e.g., melting and aqueous fluid-rock interaction, Iwamori and Albarede, 2008). One of the independent components (IC2) represents ‘anciently subducted aqueous fluid component’ stored for 300 to 900 million years in the mantle, and defines the fluid component-rich (=positive IC2) eastern hemisphere, while the western hemisphere shows the opposite polarity. We have also found a striking geometrical similarity between the IC2 and the inner core hemispheric structures (Iwamori and Nakamura, 2015): the eastern hemisphere shows positive IC2 in the mantle and high seismic velocities in the inner core. Combining these constraints, we propose ‘top-down hemispherical dynamics’: focused subduction within and around the supercontinent has created a fluid component-rich hemisphere with a lower temperature, compared to the oceanic mantle. The colder hemisphere seems to have been anchored to the asthenosphere during the continental dispersal, and may affect the temperature and growth rate of the inner core, resulting in the coupled hemispherical structures in the mantle and the core.

Keywords: mantle, core, supercontinent, hemispherical structure, isotope, independent component analysis

下部マントルに沈み込むスラブの運動と水の輸送 Slab dynamics and water transport in the lower mantle

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地表から固く冷たいスラブがマントルへ沈み込むことは、地球マントル対流を特徴づける要因のひとつであると考えられている。最近の研究では、660-km 相境界を超えたスラブは急に太くなっていることを示しているが、分解能の限界により下部マントルでのスラブの運動や変形を定量的に議論することは困難である。一方、LLSVPsの形状や親水元素の分布などにより、スラブの沈み込みとマントル最下層構造との間には関連性が示されている。本研究では、人為的な力を与えることなく対流を再現できる数値モデルを用い、下部マントルへ沈み込むスラブの運動と最下層構造のより自然的な相互作用を調べる。

2次元のマントル対流モデルによってプレートの運動と水輸送を数値シミュレーションにより解く。さらに、実験に基づく含水鉱物の相図 (Iwamori, 2004, 2007) と水の輸送を組み込む。マントル最深部には、周囲のマントルより重い層と post-perovskite (PPV) への相変化を導入する。今回は、粘性率と熱膨張率の深さプロファイル、スラブの降伏強度を変化させて計算を行う。

今回の結果では、スラブ強度が低い時は (200 MPa) 下部マントルの粘性率にかかわらずスラブは座屈する。スラブの座屈の太さや降下速度は、下部マントルの粘性率の深さ変化の影響を受ける。さらに、熱膨張率が深さ依存性を持つ時は、スラブが硬い時 (300 MPa) でも座屈は生じる。これは、熱膨張率によりマントル深部ではスラブが持つ下向きの浮力が減少する効果で説明できる。さらにこの場合、マントル浅部では、逆に海洋リソスフェアの沈み込みが促進されるために、表面プレートは下部マントルのスラブの影響を受けにくくなり、観測によるプレート運動と調和的になる。

スラブの形状の違いは下部マントルでの水が移動と最下層での高密度物質の構造に変化をもたらす。座屈したスラブは、褶曲の間に水を含むために、水は沈み込み帯前後面に広がる。一方、スラブがまっすぐに沈み込むとコア-マントル境界での接触領域の粘性が低くなり、スラブ直上に形成される含水層がスラブから剥がれるために、沈み込み帯前面に多く分布するようになる。このような水が前面に多く分布する特徴は、PPV 相のスラブの粘性率が大きく低下するときにも起こる。各物理パラメータがマントル対流に与える影響を調べた今回の研究から、下部マントルスラブ構造および CMB 近傍の物質構造には降伏強度または熱膨張率が、さらに水の分布には PPV の粘性降下が重要な役割を果たしている可能性があると考えられる。

キーワード: 自由対流, スラブの沈み込み, スラブの座屈, 最下部構造, 水輸送

Keywords: free convection, subducting slab, slab buckling, lower mantle structure, water transport

表面波の位相及び振幅を用いた北米大陸の高精度マッピング Mapping the North American continent with inter-station phase and amplitude data of surface waves

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The investigation of a three-dimensional upper mantle structure using seismic surface waves has been generally based on the measurements of phase delay. The lateral heterogeneity in the Earth, however, not only affects the phase of surface waves, but also modulates their amplitude through focusing/defocusing effects. Such amplitude anomalies caused by elastic focusing are dependent on the second derivative of phase velocity across the ray path, and thus they are sensitive to short-wavelength structure than the conventional phase data, which should be useful for improving the lateral resolution of phase velocity models. In this study, we collect a large-number of inter-station phase velocity and amplitude ratio data working with a non-linear waveform fitting technique using USArray seismograms. Phase velocity maps of North America are then constructed using both phase and amplitude data of both Rayleigh and Love waves to check the validity and utility of inter-station amplitude measurements for enhancing the quality of the phase velocity models.

The phase velocity maps derived only from phase data reflect large-scale tectonic features well; e.g., slow anomalies in the tectonically active western U.S. and fast anomalies in the eastern cratonic region. To the contrary, phase speed models derived from amplitude data tends to emphasize smaller-scale structures characterized by strong lateral velocity gradients; e.g., significant slow anomalies in Snake River Plain and Rio Grande Rift, where the local amplification due to elastic focusing has been observed at USArray stations. Our results indicate that inter-station amplitude-ratio data reflect the effects of the second derivatives of phase velocity distribution well, and are extremely useful for reconstructing shorter-wavelength elastic structures. Thus, the measurements of inter-station amplitude ratios across a dense seismic array can be used to enhance the horizontal resolution of phase velocity models of surface waves.

Keywords: surface wave, phase velocity, amplitude, tomography, North America

アセノスフェア熱イベントの検出：マントルゼノリスからのアプローチ Detection of an asthenospheric thermal event: approach from lithospheric mantle xenoliths

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Convective mantle heat flux through the continental lithosphere is not well constrained because of high heat generation in the overlying continental crust (Sclater et al., 1980; Pollack et al., 1993; Jaupart et al., 2007; Jaupart and Mareschal, 2007). Its proper estimation and its temporal variation through the earth's history is important to know the overall rate of heat loss from the convecting interior of the earth (Labrosse and Jaupart, 2007; Korenaga, 2008), although the earth is thought to be losing most of its internal heat through the oceanic lithosphere (~70%; Jaupart et al., 2007; Mareschal et al., 2012). The sub-continental lithosphere-asthenosphere boundary (LAB) is the interface through which entropy transported to the asthenosphere beneath the LAB via mantle convection from the depth of the earth is passed on to the entropy transfer in either steady or transient state through the sub-continental lithosphere (Jaupart and Mareschal, 2007; Michaut and Jaupart, 2007; transient important). There are three important mechanisms of entropy transfer through the LAB: heat conduction, solid-state flow, and magmatism (Jaupart and Mareschal, 2007). The upper most zone of the asthenosphere acts as the upper thermal boundary layer of the convecting mantle and that the heat was transferred via heat conduction in the continental lithosphere with or without LAB modification (thickening or delamination/thermal erosion of the lithosphere; Moore et al., 1999; Jurine et al., 2005). Another important aspect of the sub-continental LAB is that it roughly corresponds to a boundary where melting and segregation of melt take place either via decompressional melting in the asthenosphere or melting of the lithospheric mantle induced by the heat input or material influx. This implies that entropy can be transferred from the convecting interior to the lithosphere via magmatism involving heat release or absorption by melting, crystallization, and open-system reactions. It is important to know where magmas are generated and crystallized during its ascent to the earth's surface in the continental region in order to evaluate the role of magmatism in heat transfer through the sub-continental LAB. If a magma generated in the asthenosphere releases heat directly on the earth's surface ending as volcanic eruption and intrusion, then heat loss via magmatism is at maximum efficiency (Ogawa, 1988). Contrary to this, if the magma releases heat within the lithosphere or crust by freezing all the melt there, it heat up the host layer. In this case, the enhancement of heat loss via magmatism depends on the depth of magma freezing, though it is higher than exclusively conductive heat transfer.

In order to examine heat transfer near the sub-continental LAB, it is important to scrutinize the thermal state and its temporal and spatial variability of the mantle material near the LAB and concomitant magma formation and its subsequent magmatism. Fortunately, we have many samples from the continental lithosphere as mantle xenoliths, though xenoliths from the asthenosphere are limited. The continental lithospheric mantle has long history of its formation and modification, but we can extract not only thermal records when xenoliths were entrapped by erupted magmas (mantle geotherm) but also their temporal change before the entrapment by carefully looking at reaction processes took place responding to various thermal and chemical changes taking place in the vicinity of the LAB.

キーワード: アセノスフェア, リソスフェア, 熱イベント, マントルゼノリス, 鉱物累帯構造

Keywords: asthenosphere, lithosphere, thermal event, mantle xenolith, mineral zoning

マックスウェル粘弾性体力学方程式の従来とは別の見方に立つ定式化 An alternative formulation of the dynamics equation system of Maxwellian viscoelastic media

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岩石からなる地殻・マントルやグリーンランド・南極に存在する氷床は普通の意味では固体であるにもかかわらず、長い時間では「流動する」とみなされ、その動きを支配する法則として粘性流体に対する Navier-Stokes 方程式が採用される (例えば McKenzie, Roberts and Weiss, 1974, ほか無数)。これは物理学として正しいのであろうか?

文献により歴史を振り返ると、氷期終了後の地殻の上昇を扱った Haskell(1935) やその同僚でマントル対流を論じた Pekeris(1935) など先駆的研究を参照しながら、McKenzie はプレートテクトニクス時代のマントル対流研究の基礎づけとなるような小論 "The viscosity of the mantle(1967, Geophys. J. roy. astr. Soc. 14, p.297)" の冒頭で次のように述べている。

Creep under low stress is by diffusion and has a linear relation between stress and strain rate; it also obeys the Navier-Stokes equation. Therefore the viscosity of the mantle may be calculated from solid state theory and also from the slow deformation of the Earth.

ここでは粘性流動の物理的基礎として、固体物理や材料工学で知られているクリープ現象、即ち固体に剪断応力をかけ続けるとそれに応じたずれ変形が生じ、時間と共にそれが大きくなる現象を固体の「流れ」とみなしていることがわかる。上の文は、この応力-歪 (速度) 関係を流体力学方程式の内部応力として採用する、と言っているととれる。

即ち、方程式中の応力をせん断歪速度で表現し、その係数を粘性率と呼ぶのである。ところが、流体運動方程式で粘性力の効果は「運動量拡散」を表すことは数学的・物理的に明らかだが、この小論や固体地球科学で想定される大きな粘性率、たとえば $10^{21} \text{Pa} \cdot \text{sec}$ を使うなら、全マントル内に光速をも超える速さで運動量が拡散することを意味し、この論は物理学として受け入れられない。(Forte, 2007; Treatise on Geophysics, Vol.1, 1.23, p813 に計算あり。しかし、この事実には驚いたり、おかしいと言ったりしていない。)

上記の英語でも、日本語でも、“応力-歪 (速度) 関係” という時、どちらが原因でどちらが結果か明白でない。「固体の流動」から思い浮かぶようにクリープ現象では外から加える力が原因で、実際、粘性率を求める実験では剪断応力を与えた上でその結果生じるクリープ速度が測定されている。これに対し、流体力学方程式における応力-歪速度関係は逆で、歪速度 (速度シア) が原因となって生じる応力を方程式の粘性項で表現することが求められているのである。この混同が考え違いの原因であろう。

このような間違いを起こさない「新しい定式化」を昨年の学会で発表した。若手から碩学に至るまで、固体地球物理学者は釈然としないようであった。今回は定式化の概略を繰り返すと同時に上記観点や、弾性力と粘性力という「二つの応力が一つの物体の中で働いているという不思議」など基本的疑問を中心に説明したい。

従来と異なる見方に立った定式化の概略を箇条書きすると次の通り (昨年発表のまとめ)。

- Maxwell によるばねとダッシュポットのモデルによる構成則を基に展開された粘弾性体の力学方程式系を再検討した。応力の時間発展を表す構成方程式は物理的根拠が不明である。
- 新しい視点ではばねとダッシュポットのモデルを考えると、力の源はばねのみで、ダッシュポットの役割は粘性力ではなく、ばねの原点の位置の時間変化を記述するものとみなせる。
- この見方に立って連続体への翻訳を考えると、応力の源は弾性歪のみでよく、一方、これまでの構成則は「弾性歪が時と共に塑性歪に転化する」ということを表す、と翻訳できる。
- このことは日常経験からも明らかであり、工学 (材料力学) の教科書で、弾塑性体における Reuss の構成則 (1930) と呼ばれるものと同じである。実験的に検証されていると記される。
- それゆえ、Maxwell 構成則に代わり、この構成則を物理的基礎に立つ法則とし、連続体の運動方程式とあわせて「弾塑性体」の基礎方程式系とするのが適切である。
- 非発散の場合の基礎方程式系は、塑性変位の時間変化 (弾性変位の原点の変化) を「流れ」とみなすと、ゆっくり変化する現象については粘性流体の方程式と数学的に同等となる。したがってその解は見かけ上粘性流体の運動となるが粘性は見かけのものにすぎない。
- 発散のある圧縮性弾性体では、相変化がおこらない限り、塑性変位に転化するのは偏差成分 (deviatoric component) のみと考えられるから、単純に粘性流体と同じにはならない。
- 発散を含む一般の場合、時間的にゆっくり変わる現象に対しては、運動方程式 (力のバランス) は準弾性平衡を保つまま、塑性歪 (偏差成分) が弾性歪の偏差成分に比例して増加するような時間変化をする。

SIT35-12

会場:106

時間:5月26日 17:30-17:45

キーワード: プレート・マントル結合系力学, マントル対流, マックスウェル粘弾性体, 粘弾性体力学, 弾塑性体力学
Keywords: plate-mantle coupled system dynamics, mantle convection, Maxwellian viscoelastic medium, viscoelastic medium dynamics, elastico-plastic medium dynamics

外核最上部の密度成層した層に存在する軸対称 MAC 波の境界モード The boundary mode of axially symmetric MAC waves can exist in the stratified layer at the top of the Earth's outer core

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外核最上部には、地震学的な観測(例えば、Helfrich & Kaneshima(2010)、Kaneshima & Helfrich(2013))や理論的な推測により、軽い物質によって密度成層した層がありそうだ、ということが分かってきた。その原因としては、Helfrich(2014)では、初期地球にあった密度成層の名残とする考えが有力だとされている。この密度成層した層には、磁力(Magnetic force)・浮力(Archimedes force)・コリオリ力(Coriolis force)のバランスによって生じる MAC 波が存在する。Braginsky(1993)では、地磁気の60年変動を説明するために、その MAC 波を軸対称に限った場合の近似解が理論的に示されている。ここでの層のモデルは、外核内部との境界において密度の不連続があり、層内では浮力振動数が一定であるようなものが用いられている。(しかし、最近の地震学の観測では、このような密度の不連続はなさそうだとされている。)その解の南北方向の位相速度はアルヴェーン波速度と浮力パラメータの積となり ($c_{lat} = V_A \cdot B_u = V_A \cdot N / f$, ここで c_{lat} は南北位相速度、 V_A はアルヴェーン波速度、 B_u は浮力パラメータ、 N は層の浮力振動数、 f はコリオリパラメーター)、鉛直構造はサイン波の重ねあわせで表される。また、減衰率は磁気拡散率に比例する。南北位相速度は浮力振動数に比例するので、観測から南北位相速度が分かれば、浮力振動数が推定できる。南北波数 $l=2$ の基本モード周波数を地磁気の60年変動と合わせるようにすると、浮力振動数は地球の自転角速度の2倍ほどになる。

私たちは、Braginsky(1993)で想定されたモデルと方程式系に、密度不連続のある層境界に局在した解(ここでは境界モードと呼ぶことにする)も存在することを見出した。この解は、Braginsky(1993)の解に比べて時間スケールが小さく、もっぱら拡散によって伝播する。また、位相は層境界から離れる方向に進む。周波数はBraginsky(1993)の解のように層内の浮力振動数には依存しない。周波数と鉛直波数は層境界における密度不連続の大きさ、南北波数などに依存している。鉛直構造は境界から離れるにつれて振幅が小さくなるような波形である。密度不連続や南北波数が大きいと、波は時間的にも空間的にも減衰が大きくなる。そのため観測で見つかるには、密度不連続と層の厚さが小さいことが必要であり、南北波数が小さい波が見つかりやすいであろう。仮に、南北波数 $l=2$ の境界モードの振動数を地磁気の60年変動と合わせるようにすると、密度のとびの割合は 10^{-4} 程度になる。なお、この波は、磁気拡散率が小さいほど、時間的にも空間的にも減衰するという層内伝播する波とは逆の性質を持っている。境界モードであるので、境界面を振動させれば、励起されるかもしれない。

キーワード: MAC 波, 外核最上部, H 層

Keywords: MAC waves, the top of the Earth's outer core, H layer

高温高圧下における鉄-水反応に対する硫黄の影響 Effect of sulfur on the reaction between iron and water under high pressure and temperature

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形成期の地球内部環境下では Fe と H₂O が反応し、FeH や FeOOH 等の含水素相が出現することが報告されているが (例えば Okuchi, 1997; Ohtani et al., 2005), 本研究では更に多成分系の反応を調べるために硫黄を加え、Fe-S-H-O 系の高温高圧下における反応関係を調べた。

この系では、反応生成物として想定される金属水素化物中の水素が低圧側で試料から散逸してしまうために、実験は高エネルギー加速器研究機構 (KEK)・放射光実験施設 (PF-AR-NE1A) で行い、その場 X 線回折法によって反応相を観察することにより、試料の相転移の観察や反応生成物の同定を行った。

高温高圧発生装置には AR-NE1 に設置されたレーザー加熱型ダイヤモンドアンビルセルを用いた。出発試料には FeS の粉末を箔状に加工したものと純水を用い、レニウムガスケットに開けた試料室に封入した。圧力測定には H₂O-VII 相の状態方程式を用いた。加熱には Nd:YAG レーザーを用いた両面加熱を行い、試料の高温部からの輻射により反応温度を推定した。

本実験は 24GPa、33GPa の圧力条件、300K-1200K の温度条件で行った。その結果、これらの温度圧力範囲では反応生成物として FeS₂(Pyrite)、dhcp-FeH_x、ε-FeOOH が観察された。ε-FeOOH の安定領域は Fe と H₂O の反応で報告されている圧力よりはるかに高圧側まで存在すること、また ε-FeOOH の高温分解条件が制約できたこと、Fe-S-H 系では水素を含む FeS が報告されているが Fe-S-H-O 系では硫化物相に有意な水素が見られないこと、Fe-S-H 系では見られなかった FeS₂ が新たに出現することなどが分かった。また、発表では回収試料の SEM-EDS 分析の結果についても報告する。

キーワード: 地球核, 軽元素, 水素, 放射光

Keywords: Earth's core, light element, hydrogen, synchrotron

遷移層から下部マントルに至る圧力でのマントル鉱物の熱伝導測定 Measurement of thermal conductivity of mantle minerals at pressures of the transition zone to the lower mantle

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Knowledge of thermal diffusivity or thermal conductivity of the mantle is vital for study of the dynamics of the Earth. So far we have measured thermal diffusivity and thermal conductivity of upper mantle minerals, i.e. olivine and garnet and hydrous phases, i.e. serpentine and talc. All those data were obtained by the experiments at pressures up to 10 GPa and temperatures to 1100 K. The measurements were conducted by a pulse-heating method of one-dimensional heat flow using the Kawai-type apparatus at the Institute for study of the Earth's interior, Misasa. This current method is a predominant one for study in deep Earth's materials under pressure. It has some advantages as follows:(1) comparatively small amount of samples (2) applicable to materials with anisotropy in thermal conduction (3) simple cell assembly. Moreover, this method enables to obtain specific heat capacity under pressure.

In order to expand pressure range the cell assembly is needed to advance by reducing its dimensions. A new cell-assembly similar to our previous one is designed for a sample of 2.6 mm in diameter and 0.6 mm in thickness. This smaller cell is installed in a 14 mm edged octahedral pressure medium in 7 mm truncated anvils. This cell enables to make measurements of the thermal properties at pressures exceeding 15 GPa, which will covers the condition in the mantle transition zone. The cell will be also applied to pyroxene samples of which sizes are necessarily limited. Test measurements were made using garnet samples. The results agree well with those of the previous experiments using the larger (18-11 and 14-8) cell, and the extrapolations to zero-pressure coincide to values of other methods. Thus, the pulse heating method will be applied for thermal property measurements of wadsleyite, ringwoodite and majorite. Using large anvils (>46 mm), the method is probable to measure the thermal conductivity of MgSiO₃ perovskite (bridgmanite). However, measurements at high temperature still have somewhat problems in precision. Materials of impulse heater and external furnace should be re-considered. The precision of measurements should be improved by well-controlled machining of the cell assembly and by refining the data acquisition system.

キーワード: マントル鉱物, 熱拡散率, 熱伝導率, 高圧力, 川井型装置

Keywords: mantle minerals, thermal diffusivity, thermal conductivity, high-pressure, Kawai-type apparatus

MgO 格子原子拡散挙動の第一原理シミュレーション Lattice diffusion in MgO crystal from first principles simulation

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Rheological property is critical to understanding the mantle convection. Diffusion creep might be the dominant deformation mechanism in the Earth's lower mantle and super-Earths' mantle (e.g., Karato, 2011). Thus several experimental and theoretical studies have tried to measure lattice diffusion coefficients under pressure, which are both still technically difficult. There are two theoretical approaches to calculate self-diffusion coefficient in solids. One is based on the static lattice energy calculation and the other is based on the molecular dynamics simulation. In the former case, it is difficult to evaluate attempt frequency and in the latter case, atoms are hardly mobile in actual computation time at the Earth's lower mantle and super-Earths' mantle temperatures. These two approaches were previously applied to MgO, one of major deep mantle constituents (Ita & Cohen, 1997; Ito & Toriumi, 2007). However reported pressure dependences of the self-diffusion coefficients are contradictive with each other particularly at high pressure over 80 GPa.

In this study, we develop a new theoretical method to calculate self-diffusion coefficient in crystals with charged vacancies (Schottky pair) within the first principles framework. This method was then applied to NaCl-type MgO. We found that the calculated pressure dependences of the self-diffusion coefficients in MgO are consistent with those of Ita & Cohen (1997). Diffusion creep viscosity of MgO was then estimated using calculated diffusion coefficients. Our activation volumes are consistent with experimental values at low pressure (Van Orman et al., 2003) and decrease rapidly with increasing pressure. It suggests that super-Earths' mantle would not be quite viscous and the constant activation volume extrapolation leads to overestimation of viscosity in the deep mantle.

This method is widely applicable to other materials including bridgmanite, post-perovskite and CsCl-type MgO, which are important to analyze more realistic planetary interior dynamics.

キーワード: MgO, 格子拡散, 地球下部マントル, スーパーアースマントル, 第一原理計算
Keywords: MgO, lattice diffusion, Earth's lower mantle, super-Earths' mantle, first principles

海洋性地殻物質の沈み込みとマントル内化学的不均質構造の形成に関する数値シミュレーション
Are LLSVPs formed in the Earth's lowermost mantle by the subduction of oceanic crusts?

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We conducted a series of numerical experiments of thermo-chemical mantle convection where a subduction is preferentially induced at a continental margin, in order to verify a hypothesis that the Large Low-Shear Velocity Provinces (LLSVPs) in the Earth's lowermost mantle are formed by subduction of oceanic crust. In this study, we adopted a model of two-dimensional rectangular box of 2900km height and aspect ratio 6 with reflective boundary condition in the horizontal direction. We placed an immobile lid as a model of surface supercontinent which covers a third of the top surface. We also put a thin layer of chemically dense materials as a model of oceanic crust, which may sink into the deep mantle along with cold descending flows from the top surface.

Our calculations showed that the subducted oceanic crusts are preferentially provided under the continent when the subduction at the margin of continent is stable. However, stable subduction caused strong convection and significantly stirred the mantle under the continent. Therefore, subducted oceanic crusts were distributed almost uniformly under the continent without accumulating on the CMB. On the other hand, the cases with unstable subduction at the margin of continent showed a long-wavelength mantle convection structure which has an ascending plume along the side wall under the continent and a descending plume at the opposite side wall. The large-scale flow gathered subducted oceanic crusts under the continent and formed large piles on the CMB.

Our results suggest that the LLSVPs are hardly formed in the presence of stable plate tectonics like the current one where a stable plate motion including subduction stirs the mantle very effectively. In other words, the formation of large thermochemical piles which are equivalent to the LLSVPs should have been completed before the plate tectonics is well established, assuming that subducted oceanic crusts are the origin of LLSVPs.

キーワード: マントル対流, 数値シミュレーション, LLSVP, プレート沈み込み
Keywords: mantle convection, numerical simulation, LLSVP, plate subduction

Experimental presentation of plate subduction using paraffin wax Experimental presentation of plate subduction using paraffin wax

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Experimental approaches using analogue materials have been widely used to understand kinematic behaviors of tectonic plates. Previously molten paraffin in a tank inside a hot water bath has been used. Although tectonic plate-like behaviors, such as inclined subduction and trench migration, have been observed, the “plate” in this case was too thin to reproduce the lithospheric strength and the heat balance through the thermal boundary layer of the Earth. In order to simulate the plate and its motion as a well-developed thermo-mechanical boundary layer on top of vigorously convecting mantle, we have developed a tank apparatus and performed preliminary experiments using paraffin wax.

To control the complex heat and convection processes and for easy observation we constructed a glass tank with an inner size 120x23x4cm. The walls are constructed from double pane glass with panes separated by air gap to reduce heat loss, and reinforced with aluminum plates and bars. The paraffin was melted from below by a copper heat-sink containing 24 ceramic heating elements. To reduce heat loss to the back wall, the wall was isolated with 8cm thick foam. At the boundary layer where the paraffin wax was sticking as it cooled down we applied NiCr wire heater to the inner walls. All heating sections were controlled by variable controllers. We cooled the top layer of the wax with a cold air flow carefully controlled with thin foam plates from a vat filled with liquid Nitrogen.

The biggest challenge was the “frosting” effect especially on the front uninsulated wall that prevented the “subduction” of the forming “crust” to deeper levels. External wall temperature was 65 °C, 70 °C was measured at the boundary level by the wire heater, while the wax inside the tank was at 80 °C. Some external force was necessary to initiate a start of subduction. The maintenance of balance between the various heaters, the wall temperature, the wax temperature and the cooling rate was critical for the successful completion of the experiment.

We observed continuous subduction and clear “crust” forming with subsequent “subduction”. We can say that our experiment properly reproduces the general features of plate motion of the earth. Artificially fracturing or weakening the boundary layer and applying a vertical, downward external force were required to initiate subduction in addition to collision of the plates. The thickness of the plate was the primary parameter controlling subduction behavior and plate motion. The plate showed elastic and plastic behavior depending on its thickness and temperature. A cold and thick “plate” did not subduct even after applying an external force, and formed a stagnant lid. A hot and thinner “plate” did not show continuous subduction behavior, plate motion stopped soon after subduction was initiated, possibly because the slab pull force from the thinner partially subducted slab was too weak. Our experiment results suggest that the driving force of subduction and plate motion is slab pull, not the thermal convection of the molten paraffin or ridge push. We will present photos and videos of the observed processes.

Improvements to the tank and heating elements design are necessary to provide better and easier control over the experiments.

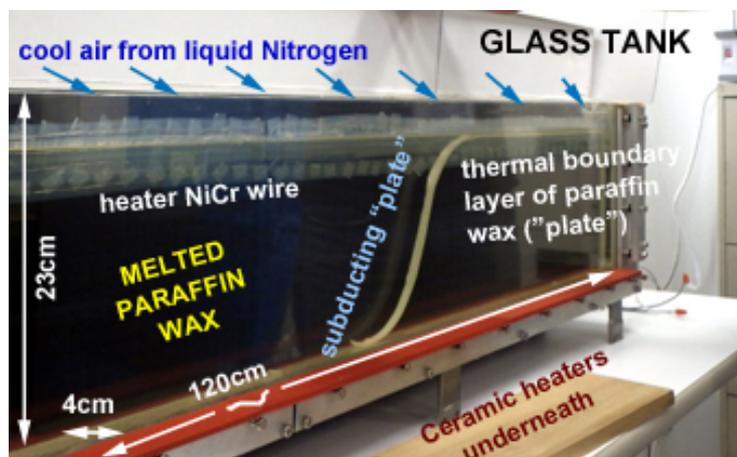
キーワード: analogue experiment, plate subduction, paraffin wax, glass walls tank, slab pull

Keywords: analogue experiment, plate subduction, paraffin wax, glass walls tank, slab pull

SIT35-P06

会場:コンベンションホール

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スーパー地球のマントル対流シミュレーション: 熱伝導率の深さ依存性と断熱圧縮の効果

Numerical experiments on mantle convection of super-Earths with variable thermal conductivity and adiabatic compression

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Recently, many extra-solar planets have been discovered by improved observation technologies. Some of these planets, called super-Earths, have small masses (up to 17 times the Earth's) and high mean density ($>5000 \text{ kg/m}^3$). Numerical modeling of mantle convection of super-Earths plays an important role in studying the occurrence of plate tectonics and the surface environments on these planets. On the other hand, when considering mantle convection of super-Earths, it is also important to take into account the difference in (hydrostatic) pressure in the mantles. Since super-Earths have high inner pressure, there must exist a strong change in physical properties and the effect of adiabatic compression. While the effects of physical properties have been intensively studied so far, those of adiabatic compression have not been well studied in the previous models of mantle convection of super-Earths. Here we conduct numerical experiments of thermal convection of highly compressible fluid in a two-dimensional rectangular box whose thermal expansivity and conductivity are dependent on depth, viscosity is dependent on temperature, in order to elucidate the mantle convection on super-Earths.

Our numerical experiments showed the change in convecting flow patterns depending on the temperature-dependence in viscosity, regardless of the depth-dependence in thermal conductivity. When a viscosity is sufficiently dependent on temperature, horizontal flow becomes dominant in the mantle, with a very weak activity of hot plumes from the base of the mantle. This flow pattern is quite similar to the "stratosphere" in the field of meteorology. In addition, we found that the occurrence of "stratosphere" is enhanced for a strong depth-dependent thermal conductivity. One reason for this is that high conductivity at depth significantly reduces the difference in temperature between the basal thermal boundary layer and isothermal core. Our study therefore suggests that the depth-dependent thermal conductivity is one of the most important agents which control the mantle dynamics of super-Earths.

キーワード: スーパー地球, マントル対流, 断熱圧縮, 熱膨張率, 熱伝導率, 粘性率

Keywords: super-Earths, mantle convection, adiabatic compression, thermal expansivity, thermal conductivity, viscosity

コア形成シミュレーションにおけるストークス流の移流陰解法 Implicit solution of the material transport of the core formation simulation

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In order to investigate the long time-scales of the global core formation process in a growing planet, we are developing the Stokes flow simulation code using MIC based techniques for material transport with a free-surface treatment. We are interested in the dynamical change of the internal structure after solidification of magma ponds/oceans during the core formation under a self-gravitating field, especially because it might lead to an initial heterogeneous structure in the deep mantle.

However current numerical solution method is difficult to solve the system coupled with the energy equation because the numerical system becomes stiff when the dynamical balancing time scale for the increasing/decreasing load by surface deformation is very short compared with the time scale associated with thermal convection. Any explicit time integration scheme will require very small time steps; otherwise, serious numerical oscillation (spurious solutions) will occur.

In this work, we propose to treat the advection as a coordinate nonlinearity, coupled to the momentum equation, thereby defining a fully implicit time integration scheme suitable for stiff problems [Furuichi and May, *Compt. Phys. Commun.* 2015]. We utilize a Jacobian free Newton Krylov (JFNK) based Newton framework to solve the resulting nonlinear equations. We also investigate efficient solution strategies to reduce the computational cost to evaluate the nonlinearity on MIC advection.

These implicit methods are implemented within FD framework [Gerya and Yuen, 2003]. We examine the solution quality and efficiency of these methods by performing numerical experiments we have performed a series of numerical experiments which clarify the accuracy of solutions and trade-off between the computational cost associated with the nonlinear solver and time step size.

キーワード: コア形成, 陰解法, 非線形ソルバ, 自由境界表面

Keywords: core formation, Stokes flow, free surface, implicit time integration, JFNK

Hi-net で計測される SKS splitting parameter のモデリング Modeling of SKS splitting parameters measured in Japan with Hi-net

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To systematically investigate the spatial variation of seismic anisotropy around Japanese islands, we measured splitting parameters (fast polarization direction ϕ , delay time δt) of teleseismic SKS phases observed by Hi-net (Ogawa et al., 2014, SSJ). The results indicated regional scale variations of splitting parameters that are apparently related to subduction systems. In order to investigate detailed anisotropic structures (fabric in mantle wedge, subducting slab, and asthenosphere), we conducted forward modeling using synthetic seismograms. We modeled the SKS phases by the ray theory. We assumed that the SKS ray is straight and that each region has homogenous anisotropy. We rigorously calculated the phase velocity in each region by solving the Christoffel matrix. The preliminary analysis indicates that the measured splitting parameters appear to be primarily affected by the A-type fabric in subducting slab (oceanic lithosphere) whose a-axis aligns in the direction of the fast axis observed at the surface by using our OBS data.

キーワード: 地震波異方性, モデリング

Keywords: seismic anisotropy, s-wave splitting, modeling

マルチモード表面波による上部マントル鉛直異方性の制約 Constraining radial anisotropy in the upper mantle with multi-mode surface waves

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The radial anisotropy of shear waves represents the differences in the propagation speeds between vertically polarized shear waves (SV) and the horizontally polarized ones (SH), and can be a key to the understanding of the dynamic processes in the upper mantle. Seismic surface waves are the most powerful tool to determine the spatial distribution of the radial anisotropy. Some recent studies have revealed the existence of a layer with strong radial anisotropy (with SH > SV) beneath the lithosphere; e.g., under the Pacific plate (Nettles & Dziewonski, 2008, JGR) and the Australian continent (Yoshizawa, 2014, PEPI). This is, however, not always the case and there are also some studies on radial anisotropy that do not show such a clear layer with SH > SV beneath the lithosphere. These differences may be related to the differences in model parameterization.

For the inversions of multi-mode phase speeds of Rayleigh and Love waves for radial anisotropy of shear waves, we can use either set of model parameters for the representation of the anisotropic S velocity; i.e., (A) SV velocity (V_{sv}) and SH velocity (V_{sh}), or (B) SV velocity (V_{sv}) and radial anisotropic parameter $\xi = (V_{sh}/V_{sv})^2$. The choice of model parameters for inversion is arbitrary, but, through synthetic experiments, we have confirmed that this difference causes non-negligible effects on the reconstruction of radial anisotropic properties of shear waves. This is mainly caused by the differences in the sensitivity kernels of Love-wave phase speeds to V_{sv} , V_{sh} and ξ .

For the set of parameters (B) [V_{sv} , ξ], Love waves always have the largest sensitivity to V_{sv} with suppressed sensitivity to ξ , and the kernel shapes for both V_{sv} and ξ are nearly identical. On the other hand, for the parameterization with (A) [V_{sv} , V_{sh}], Love wave phase speeds are controlled primarily by the kernels for V_{sh} , which have the largest sensitivity to Love wave phase speeds with little influence from V_{sv} , which can be better (and independently) constrained by Rayleigh waves.

Such intrinsic differences in the sensitivities of surface waves can lead to the different results in the estimation of radial anisotropy. Our synthetic experiments suggest that the parameterization with [V_{sv} , V_{sh}] would be preferable particularly when the radial anisotropy with SH > SV is caused by anomalously slow SV velocity, which is consistent with the recent anisotropy models reported in the fast moving Pacific and Australian plates. We have also found that the strong dependence of the retrieved anisotropy on the initial model, when we use [V_{sv} , ξ] as model parameters.

Keywords: radial anisotropy, surface waves, upper mantle, lithosphere, asthenosphere

日本列島下の遠地地震P波・S波トモグラフィー Teleseismic P- and S-wave tomography beneath Japan Islands

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日本列島は、4つのプレートが隣接し相互に作用しあう複雑なテクトニクスを有しているため、その地下構造は非常に不均質である。これまでに近地地震データを用いた数多くの地震波トモグラフィ研究により、詳細な日本列島下の3次元速度構造の推定がなされてきた。しかしながら、その研究対象領域は地殻や上部マントルといった浅部に限定されており、日本列島深部(深さ 200-700 km)の構造を推定した研究は少なく、その構造はあまり分かっていない。日本列島深部の構造を調べることは、沈み込むスラブやマントル上昇流の深部での様子や相互の関連などを理解できると考えられ、マントルダイナミクス解明において非常に重要である。

本研究では、遠地地震の相対走時残差を使用した遠地地震トモグラフィにより日本列島深部のP波、S波の速度構造を推定した。P波を用いた同様の研究は過去にも行われてきたが(例えば, Zhao et al., 1994, 2012; Abdelwahed and Zhao, 2007 など), S波を用いた研究は少ない。P波とS波両方のデータを用いることで、速度構造だけでなくポアソン比の分布も得られ、日本列島地下深部の構造をより詳細に把握することができる。

本研究で用いた遠地地震の相対走時残差は、P波においては、先行研究(Zhao et al., 1994, 2012; Abdelwahed and Zhao, 2007)により集められた日本からの震央距離 30~100° の地震 360 個の中から、日本全土で観測されており、かつ 100 点以上の観測点で観測された 130 個の遠地地震と、震源の分布がなるべく均等になるように新たに読み取った 38 個地震の計 168 個の遠地地震より約 6 万個のP波走時データである。S波においては、震源の分布が均等になるように選んだ 56 個の遠地地震より約 4 万個のS波データを使用した。また、浅い速度構造を補正するために、気象庁の一元化カタログから 100 点以上で観測された近地地震約 3 千個を用いた。解析手法には、近地地震と遠地地震のデータを同時にインバージョンすることができる Zhao et al. (1994, 2012) のトモグラフィ計算プログラムを用いた。

得られた結果より、P波、S波ともに背弧側から続く低速度異常体が見られ、これはマントルウェッジ内のコーナーフローであると考えられる。関東地方以西においては、内陸下の稍深発地震の震源に沿った高速度異常を見ることができ、沈み込むフィリピン海スラブであると考えられる。特に九州地方になるとこの高速度異常は深さ 400 km 程にまで達する様子がP波、S波の両方で見ることができ、フィリピン海スラブは九州下ではマントル遷移層まで沈み込んでいると考えられる。発表ではこれらに加え、ポアソン比やマントル深部物質の物性に関するR値($\ln V_s / \ln V_p$)を求めることで、日本列島の深部構造をより詳細に議論したい。

Cape Verde ホットスポットとマントルプルーム Seismic evidence for a mantle plume beneath the Cape Verde hotspot

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The Cape Verde hotspot is located in the African plate, about 2000 km east of the nearest plate boundary. It is composed of a group of late Cenozoic oceanic islands resting on a broad bathymetric swell on mature (>110 Ma) oceanic lithosphere. This hotspot has a positive surface heat flow, high geoid anomaly, and long-term volcanism. The last known volcanic eruption occurred at Fogo volcano in 1995.

We determined P- and S-wave tomography of the upper mantle beneath the Cape Verde hotspot using arrival-time data measured precisely from three-component seismograms of 106 distant earthquakes recorded by a local seismic network. Our results show a prominent low-velocity anomaly imaged as a continuous column <100 km wide from the uppermost mantle down to about 500 km beneath Cape Verde, especially below the Fogo active volcano, which erupted in 1995. The low-velocity anomaly may reflect a hot mantle plume feeding the Cape Verde hotspot.

キーワード: Cape Verde, ホットスポット, マントルプルーム, 地震波トモグラフィ, マントル遷移層
Keywords: Cape Verde, hotspot, mantle plume, seismic tomography, mantle transition zone